

# Climate-Change Simulations with HSPF 5.3.0 Model of the Chesapeake Bay Watershed

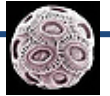


*Maria Herrmann  
and  
Ray Najjar*

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# Collaborators

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Michael Barnes, University of Maryland, Baltimore County

Mark R. Bennett, U.S. Geological Survey

Lauren E. Hay, U.S. Geological Survey

Lewis Linker, U. S. Environmental Protection Agency

Christopher R. Pyke, U.S. Green Building Council

Kevin G. Sellner, Chesapeake Research Consortium

Gary Shenk, U. S. Environmental Protection Agency

Denice H. Wardrop, The Pennsylvania State University

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To assess how future climate change will affect streamflow, nutrients, and sediment in the Chesapeake Bay watershed

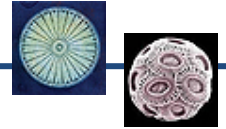


- The Chesapeake Bay is profoundly affected by its watershed, which contains more than 100,000 streams and rivers and whose area is about 14 times larger than the area of the estuary (USEPA, 2010)
- Changes in air temperature and precipitation associated with climate change are likely to alter inputs of fresh water, sediment, and nutrients from the watershed to the bay
- Climate-induced changes in streamflow and loads of nutrients and sediments have been identified as some of the largest uncertainties in the Chesapeake Bay's response to climate forcing (Najjar et al., 2010).



- Watershed fluxes were simulated using HSPF\* 5.3.0 model of the Chesapeake Bay watershed (Bicknell et al., 1997; USEPA, 2010 )
- Output of six General Circulation Models (GCMs) was used to create forcing for the climate-change runs of the hydrological model (Hay et al., 2011)
  - A2 emissions scenario (high-end, atmospheric CO<sub>2</sub>~750 ppm by 2090)
  - only changes in the mean annual cycles of T and P; no changes in variability
- 10-year simulations of the hydrological model
  - baseline: 1990 – 1999
  - climate-change runs: 2086 – 2095
- Change ( $\Delta$ ) = climate run – base run
- Watershed-wide results

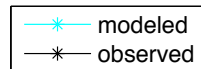
\* Hydrological Simulation Program in Fortran



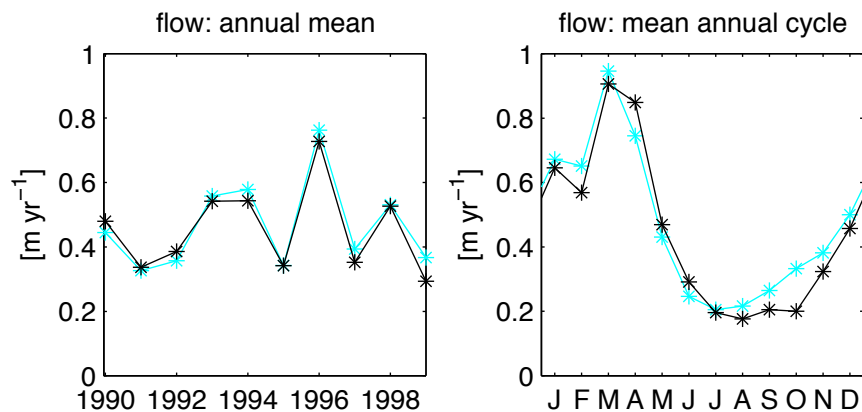
	GCM name	Originating group
1	BCCR-BCM2.0	Bjerknes Centre for Climate Research, Norway
2	CSIRO-Mk3.0	Commonwealth Scientific and Industrial Research Organization, Australia
3	CSIRO-Mk3.5	Commonwealth Scientific and Industrial Research Organization, Australia
4	INM-CM3.0	Institute for Numerical Mathematics, Russia
5	MIROC3.2(medres)	National Institute for Environmental Studies, Japan
6	CCSM3	National Center for Atmospheric Research, USA

- Monthly mean T and P from the multi-model dataset archive of the World Climate Research Programme Coupled Model Intercomparison Project, phase 3 (CMIP3)  
[http://www-pcmdi.llnl.gov/ipcc/model\\_documentation/ipcc\\_model\\_documentation.php](http://www-pcmdi.llnl.gov/ipcc/model_documentation/ipcc_model_documentation.php)

# Hydrological model evaluation



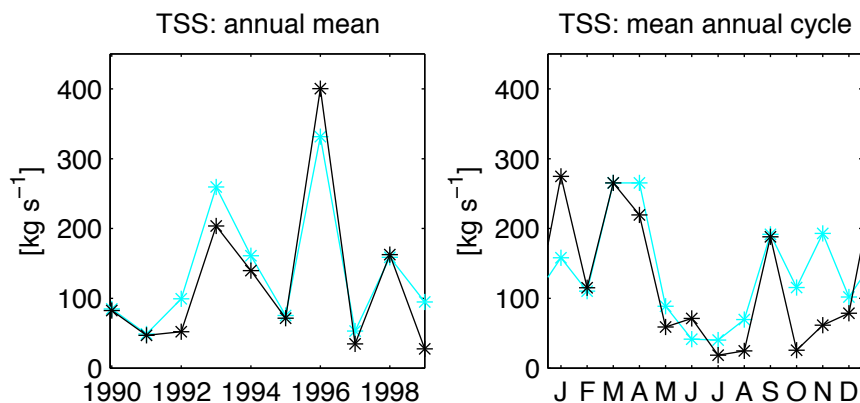
## A. Streamflow (Q)



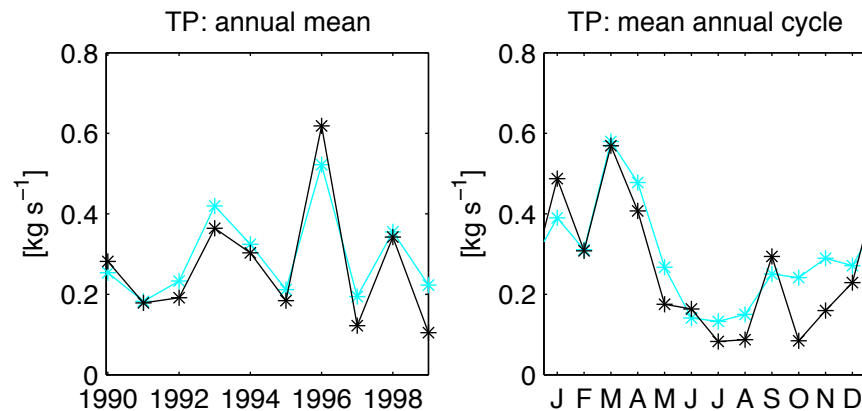
## C. Total nitrogen (TN)



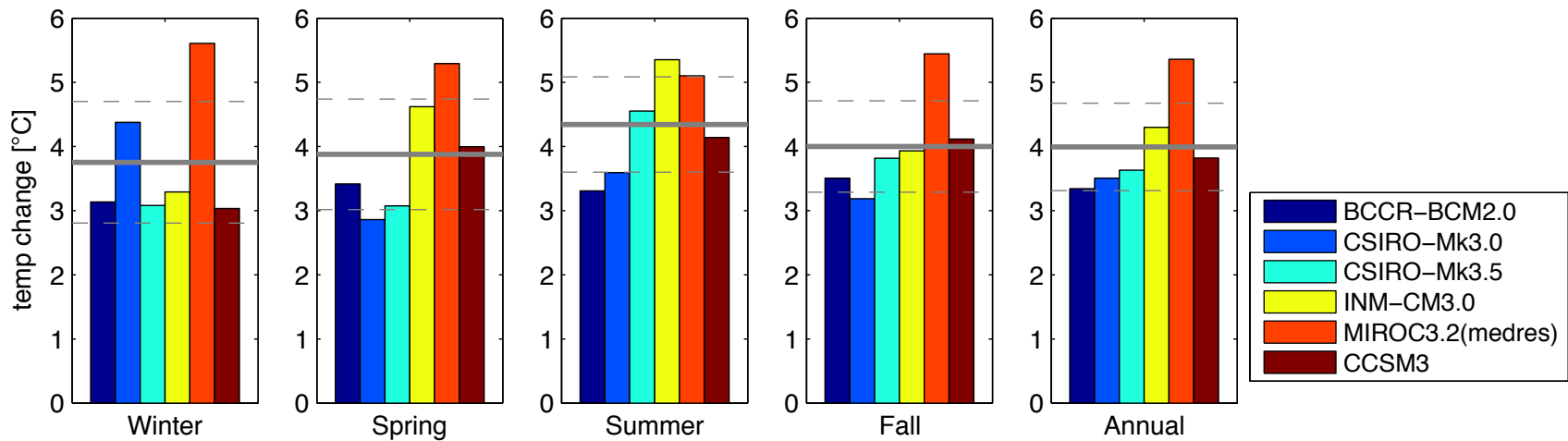
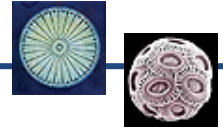
## B. Sediment (TSS)



## D. Total phosphorus (TP)



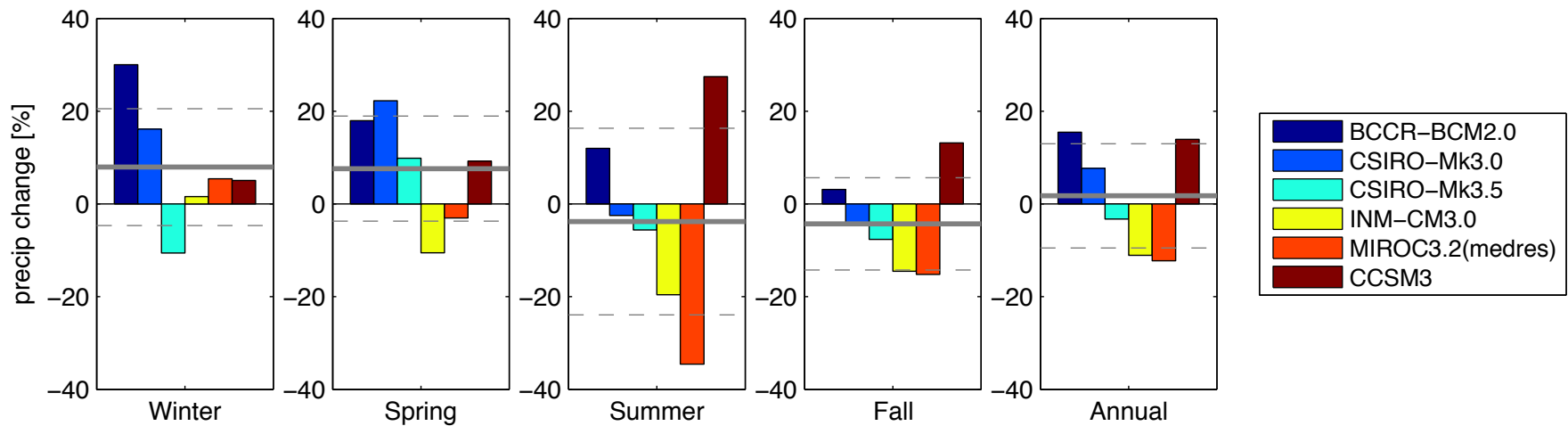
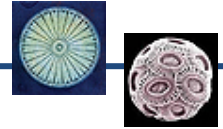
# Forcing: atmospheric temperature (T)



Annual summary [°C]	MEAN	STD	MIN	MAX
Baseline average	11.0	0.8	10.2	12.3
Projected changes	4.0 (36 %)	0.8 (7 %)	3.3 (30 %)	5.4 (49 %)

- All GCMs project warming in every season
- Largest warming in the summer
- Projected warming is far outside the bounds of interannual variability

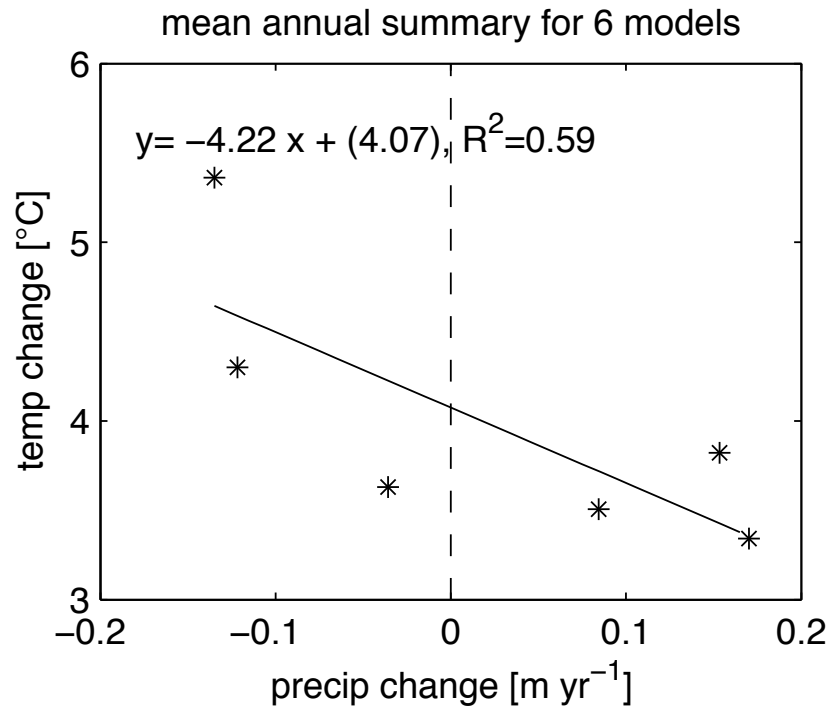
# Forcing: precipitation (P)



Annual summary [m yr <sup>-1</sup> ]	MEAN	STD	MIN	MAX
Baseline average	1.1	0.2	0.9	1.4
Projected changes	+0.02 (2 %)	0.14 (12 %)	-0.13 (12%)	+0.17 (15 %)

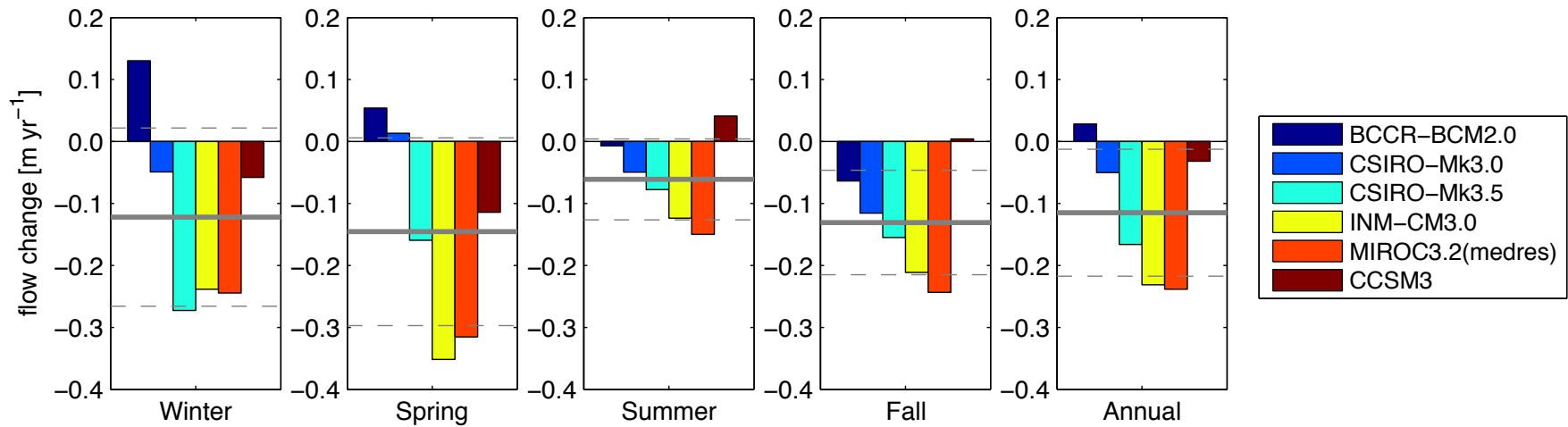
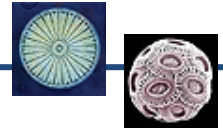
- Considerable variability in P projections
- Winter and spring P increases, on average
- Summer and fall P decreases, on average
- Projected changes are within the bounds of interannual variability

# Summary of projected changes in forcing



- Projected annual changes in T and P for the six models are negatively correlated, so models that get drier tend to warm more than models that get wetter
- Lower soil moisture due to lower precipitation would lead to a decreased latent cooling and, thus, more warming for a given amount of incident solar radiation

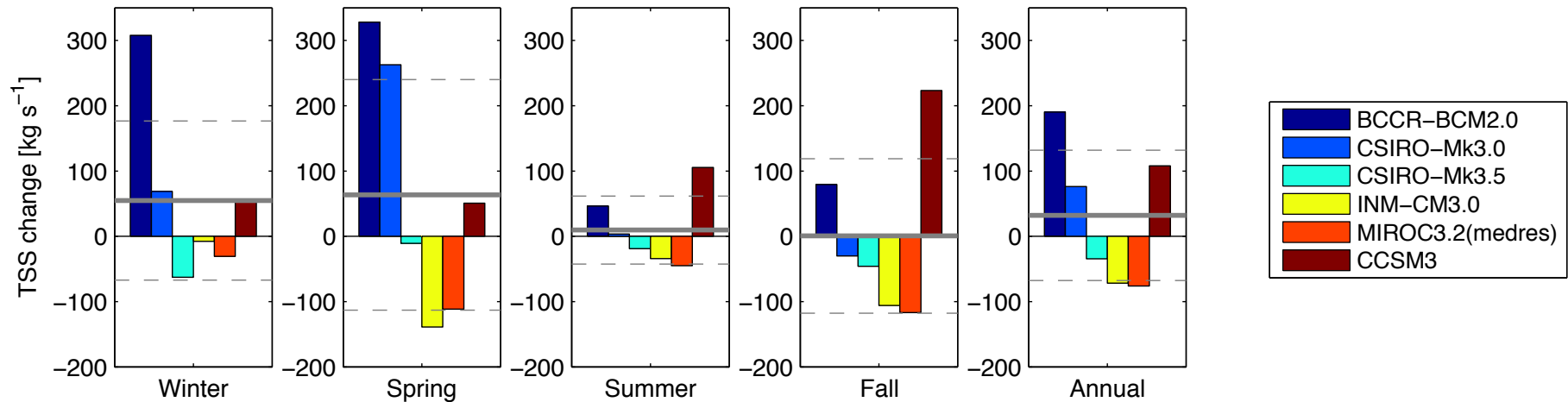
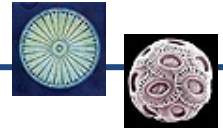
# Simulated Q response



Annual summary [m yr <sup>-1</sup> ]	MEAN	STD	MIN	MAX
Baseline average	0.5	0.1	0.3	0.8
Projected changes	-0.11 (25 %)	0.11 (25 %)	-0.24 (51%)	+0.03 (6 %)

- Dramatic decline in all seasons
  - 5 out of 6 runs show a decrease in winter, summer and fall
  - 4 out of 6 runs show a decrease in spring
- Summer decline is the least pronounced
- Projected changes are comparable to interannual variability

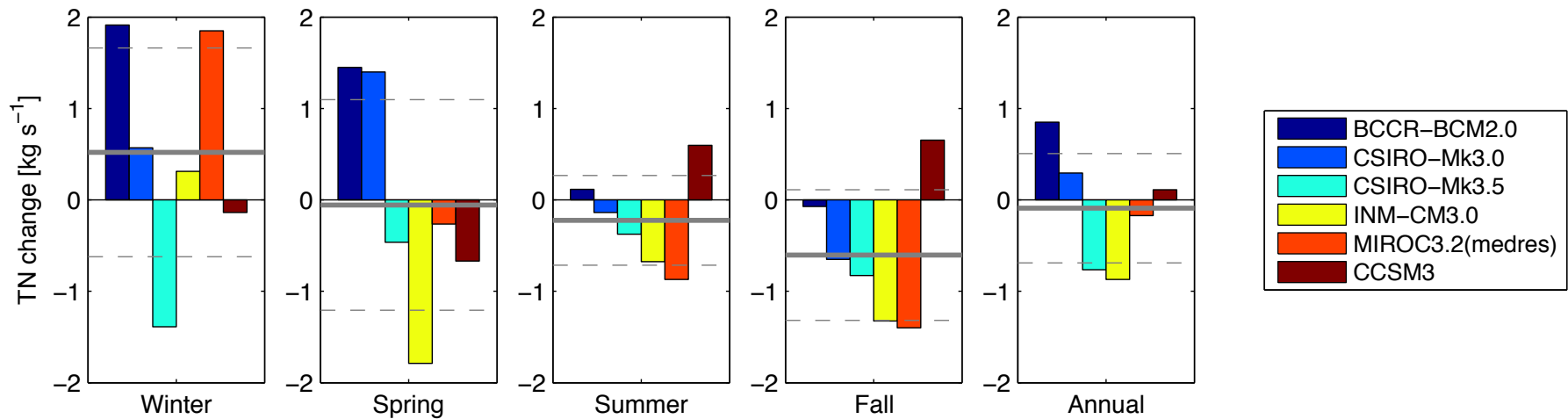
# Simulated TSS response



Annual summary [ $\text{m yr}^{-1}$ ]	MEAN	STD	MIN	MAX
Baseline average	137	93	49	331
Projected changes	+32 (23 %)	109 (80 %)	-76 (56%)	+190 (139 %)

- Projected changes are highly variable
- Winter and spring fluxes increase, on average
- Summer and fall fluxes remain essentially unchanged, on average
- Projected changes are comparable to interannual variability

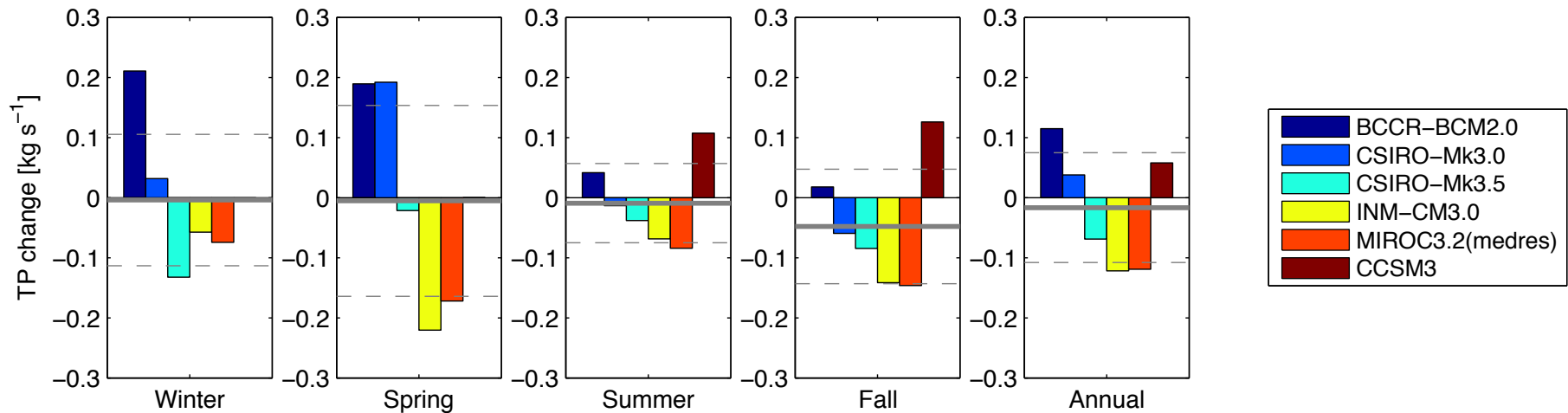
# Simulated TN response



Annual summary [ $\text{m yr}^{-1}$ ]	MEAN	STD	MIN	MAX
Baseline average	4.6	1.1	3.4	6.8
Projected changes	-0.1 (2 %)	0.7 (14 %)	-0.9 (20 %)	+0.9 (20 %)

- Projected changes are highly variable
- Winter fluxes increase, on average
- Spring, summer and fall fluxes decrease, on average
- Projected changes are comparable to interannual variability

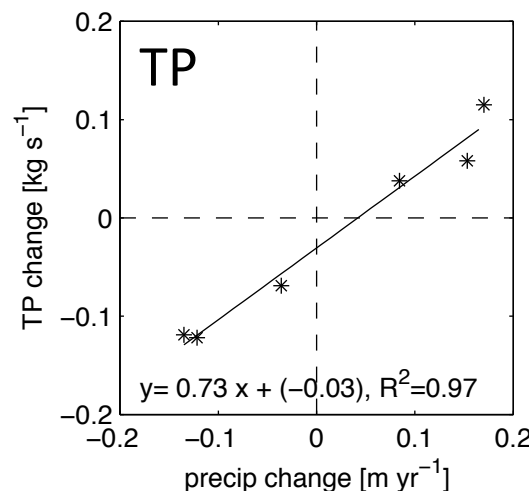
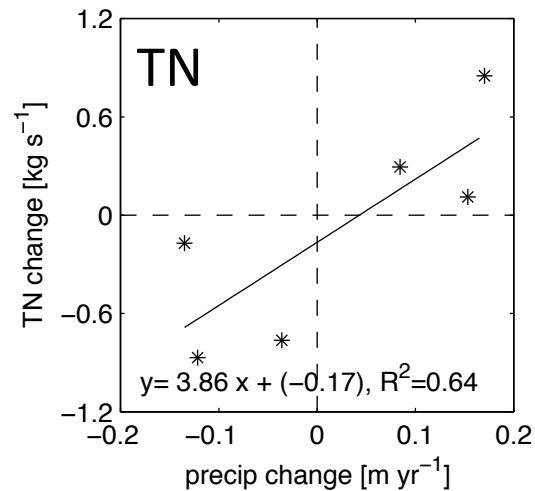
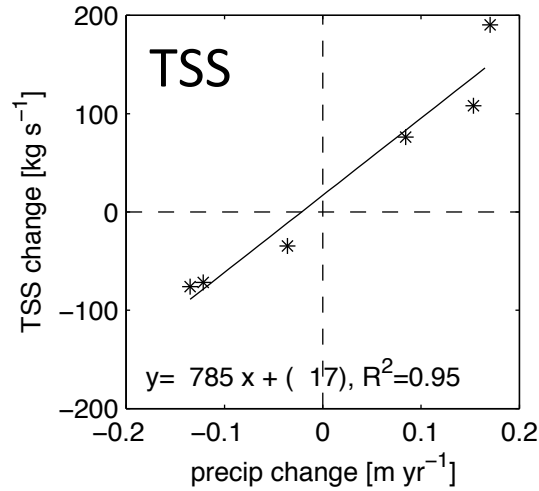
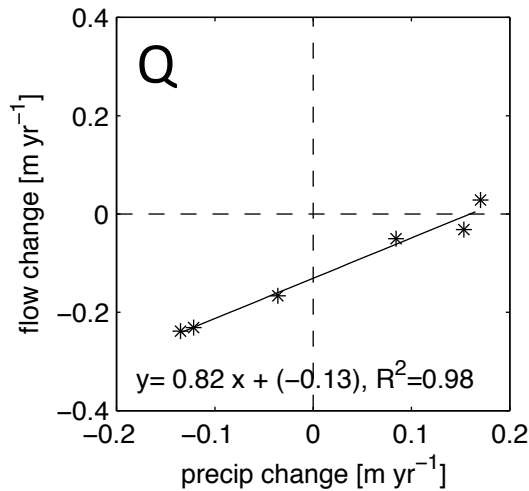
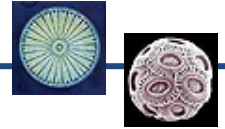
# Simulated TP response



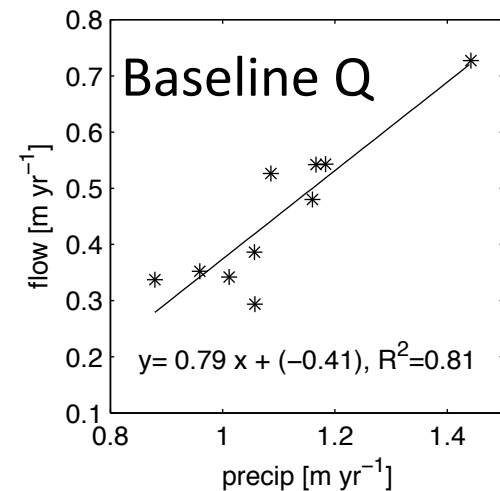
Annual summary [ $\text{m yr}^{-1}$ ]	MEAN	STD	MIN	MAX
Baseline average	0.29	0.11	0.18	0.52
Projected changes	-0.02 (6 %)	0.10 (34 %)	-0.12 (40 %)	+0.12 (40 %)

- Projected changes are highly variable
- Winter and spring fluxes remain essentially unchanged, on average
- Summer and fall fluxes decrease, on average
- Projected changes are comparable to interannual variability

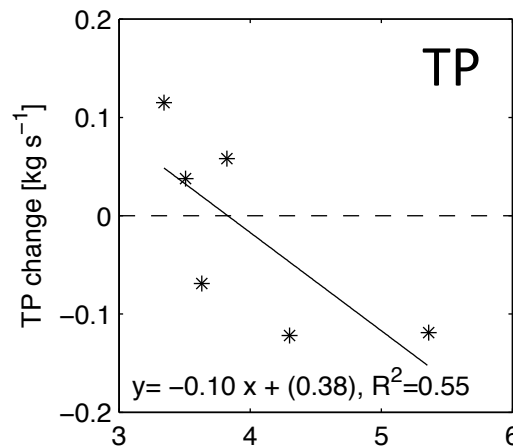
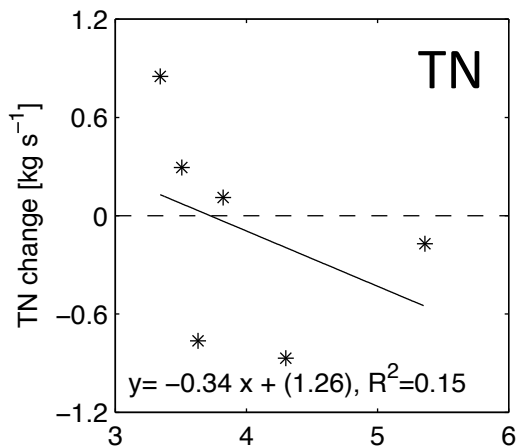
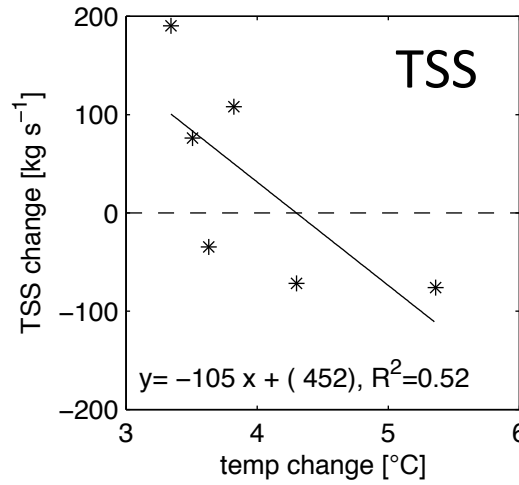
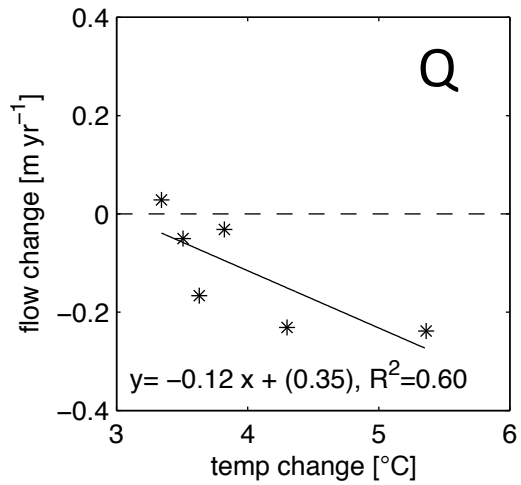
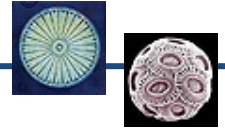
# Sensitivity to P



- Positive correlation with P
- P accounts for a considerable fraction of variability
- The y-intercepts can be interpreted as the projected flux changes due to T alone
- The slope of the Q vs. P is similar to what was found for the baseline interannual variability (slope = 0.79,  $R^2 = 0.81$ )

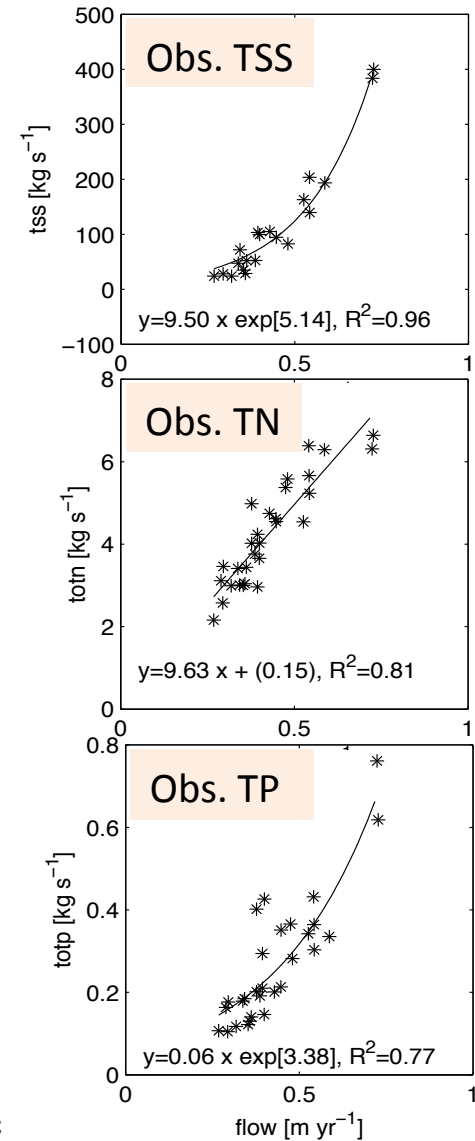
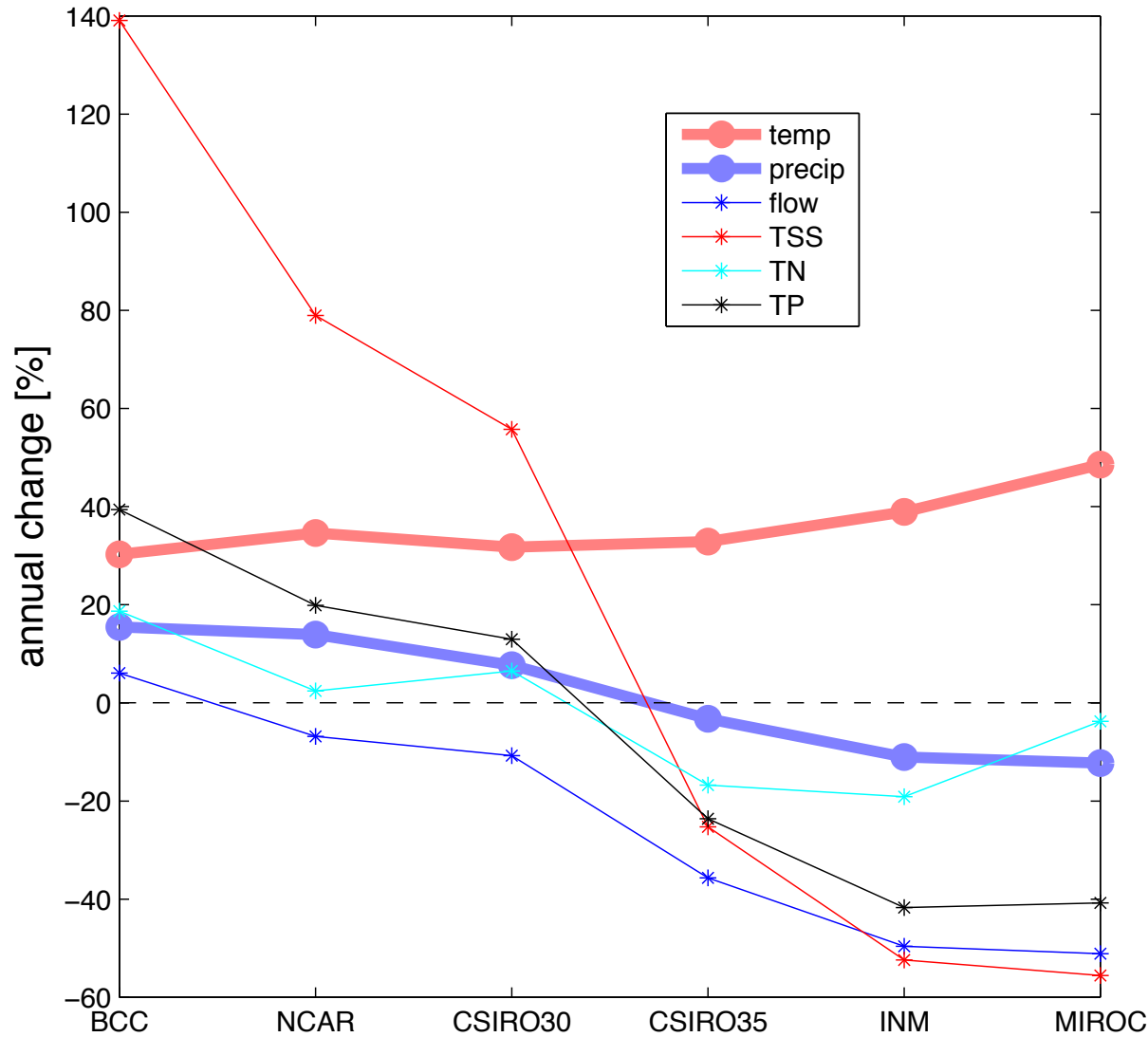
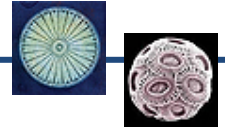


# Sensitivity to T

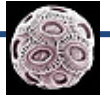


- Negative correlation with T
- T accounts for a smaller fraction of variability compared with P
- For TN, the relationship is particularly weak
- For Q, have the strongest relationship

# Summary of annual response to climate forcing



# Comparison to other Q projections



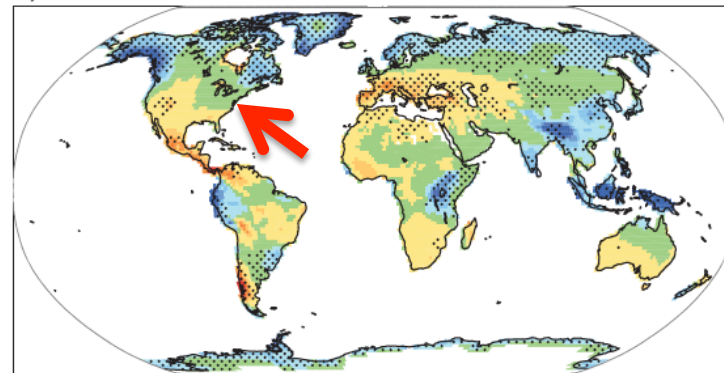
## Najjar et al. 2009

- End of 21<sup>st</sup> century
- Mid-Atlantic region
- Coupled models
- Large variability  $\Delta Q$

Reference	Region	Annual streamflow change (%)
McCabe and Ayers (1989)	Delaware River Basin	-39 to 9
Moore et al. (1997)	Mid-Atlantic/New England	-32 to 6
Najjar (1999)	Susquehanna River Basin	24 $\pm$ 13
Wolock and McCabe (1999)	Mid-Atlantic	-25 to 33
Neff et al. (2000)	Susquehanna River Basin	-4 to 24
Frei et al. (2002)	Southeastern New York	-28 to 10
Hayhoe et al. (2007)	Pennsylvania and New Jersey	9 to 18

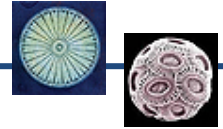
## Meehl et al. 2007, IPCC 4AR

- End of 21<sup>st</sup> century
- Coupled models
- 15-model mean projections
- $\Delta Q = +0.1 \text{ mm d}^{-1} = 0.04 \text{ m yr}^{-1}$



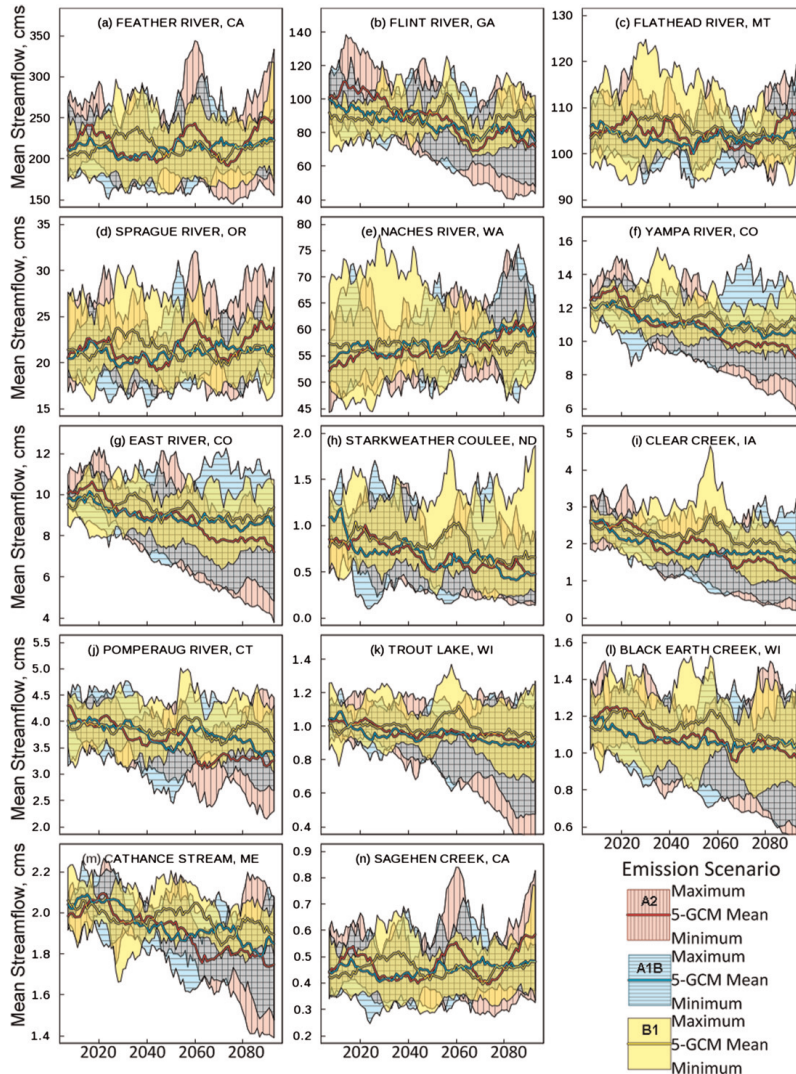
Annual Q summary [m yr <sup>-1</sup> ]	MEAN	STD	MIN	MAX
Baseline average	0.5	0.1	0.3	0.8
HSPF projections	-0.11 (25 %)	0.11 (25 %)	-0.24 (51%)	+0.03 (6 %)

# Comparison to other Q projections (contd.)



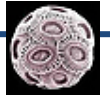
Hay et al. 2011

End of 21<sup>st</sup> century projections for 14 U.S. watersheds: hydrological model

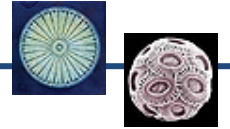


Milly and Dunne 2011

- used Hay et al. 2011 hydrological model projections
- compared to climate – land surface coupled models (Coupled Model Intercomparison Project, CMIP)
- $\Delta$ PET according to T-based empirical formulation was typically 3 times larger than that of the explicit calculation of the surface energy balance
- uncoupled simulations may lead to unrealistically large flow reductions: empirical PET formulations calibrated in the present climate might cause an overestimation of ET when used for future climate conditions

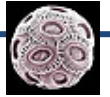


- $\Delta T = +4\text{ }^{\circ}\text{C}$  → outside the bounds of interannual variability
- $\Delta P$  range from -12 % to +15 %
- Annual  $\Delta T$  and  $\Delta P$  for the 6 GCMs are negatively correlated
- Dramatic decline in streamflow: annual  $\Delta Q = -25\%$
- $\Delta TSS$  range from -56% to + 139%
- $\Delta TN$  range from -20% to + 20%
- $\Delta TP$  range from -40% to + 40%
- All parameters are strongly correlated with P
- All parameters are weakly anti-correlated with T
- Q has the strongest anti-correlation with T: real or due to PET parameterization?



- What are the separate effects of PET, T, and P?
  - T change only (6 runs)
  - P change only (6 runs)
  - PET change only (6 runs)
  
- Are HSPF-based projections flows comparable to those from coupled models?
  - use best regional hydrological model predictions available for our study region from coupled climate-land surface models: mid-century (2041-2070) projections are available from the North American Regional Climate Change Assessment Program (NARCCAP)
  - for comparison with Milly and Dunne (2011), also use Couple Model Intercomparison Project (CMIP) models
  - 6 additional HSPF simulations that overlap with the NARCCAP period

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